

APPENDIX B

Experimental Procedures For Sedimentological And Mineralogical Analyses

B.1 Textural analysis

The core samples were subjected to grain size analysis. In the sand layer, about 60-80g. sample was taken. This was subjected to sequential 10% HCl and 30% v/v hydrogen peroxide treatment to remove the carbonates and organics. The sample was then wet sieved through 62.5 μ sieve (BSS No. 240). The finer than 62.5 μ was used to determine the amount of silt and clay using standard pipette analysis method (Carver, 1971) employing the principle of Stoke's settling. The sand-silt-clay percentages were then computed. Results are given in Table B.1

B.2 Heavy mineral separation

Heavy minerals were separated from the sand fraction of the sample using bromoform ($\rho=2.85$). The samples were transferred to a separating funnel which had been filled with bromoform. After a few hours time, the heavy mineral fraction was collected in a beaker and washed. The remaining bromoform was kept aside for further use. The heavy minerals were mounted on glass slides and identified under a petrological microscope.

B.3 X- Ray diffraction

This was used for the identification of clay minerals to determine the provenance for Horizon-1 and Horizon-3. Calcium carbonate and organic matter were removed by dilute acetic acid and 6% v/v hydrogen peroxide respectively. Clay fraction ($<2\mu$) was collected using Stoke's settling principle and concentrated by centrifugation. It was then pipetted onto glass slides and dried. The slides were scanned between 3° to 30° (2θ) at

1°/min/cm using a CuK_α source. The clay minerals were identified from the charts using the powder diffraction data file, published by International Centre for Diffraction Data, Pennsylvania.

Smectite has a broad peak in the region, 7.3° to 5.9° (2 θ). However the peak for vermiculite also appears in this region (6.22°). The presence of smectite was confirmed using glycolation method. To the natural clay sample, 1N MgCl_2 was added and left overnight. The next day this mixture was washed with distilled water. A slide was subsequently prepared from this clay sample. This slide was then kept in a dessicator with ethyl glycol at 60°C for 6 hours. If the peak in the natural slide is of smectite and not vermiculite, the treated slide will show the widening of the peak.

Also both kaolinite and chlorite have the same peak at 12.4° (2 θ). Differentiation of kaolinite and chlorite was carried out using the thermal treatment. To the natural slide was added a 1N solution of KCl and left overnight. The next day this mixture was washed with double distilled water and a slide prepared. This slide was then heated to 550°C. The peak remains if the mineral is chlorite.

For the purpose of semi-quantification, the area under the peak was graphically computed. Using the following formula (Biscaye, 1965), relative clay mineral percentages calculated,

$$\text{smectite} + 4 * (\text{illite}) + 2 * (\text{kaolinite}) + 2 * (\text{chlorite}) = 100\%$$

The results are given in Table B2.

Table B.1 Results of grain size analyses* on samples from Nal Sarovar core.

| Lab no. | Depth (cm) | Sand (%) | Silt (5) | Clay (%) |
|---------|------------|----------|----------|----------|
| N-4 | 10-20 | 3.84 | 31.19 | 64.97 |
| N-34 | 86.5-90 | 3.38 | 14.20 | 82.42 |
| N-75 | 195-207 | 7.33 | 31.78 | 60.89 |
| N-83 | 280-290 | 14.41 | 45.47 | 40.42 |
| N-99 | 390-393 | 85.98 | 10.9 | 3.12 |
| N-100 | 393-417 | 64.62 | 27.53 | 7.85 |
| N-105 | 477-522 | 91.34 | 6.95 | 1.71 |
| N-107 | 540-552 | 82.79 | 10.31 | 6.90 |
| N-113 | 600-610 | 64.43 | 24.44 | 11.13 |
| N-118 | 685-702 | 70.04 | 13.72 | 16.24 |
| N-124 | 775-790 | 44.76 | 37.62 | 17.62 |
| N-126 | 887-907 | 65.90 | 26.60 | 7.50 |
| N-127 | 907-920 | 71.59 | 19.32 | 9.09 |
| N-132 | 987-1000 | 73.49 | 22.71 | 3.82 |
| N-136 | 1065-1102 | 51.15 | 39.03 | 9.82 |
| N-143 | 1217-1225 | 81.07 | 10.81 | 8.12 |
| N-147 | 1285-1299 | 37.26 | 49.87 | 12.87 |
| N-148 | 1299-1324 | 41.96 | 43.95 | 14.09 |
| N-153 | 1405-1430 | 80.46 | 15.27 | 4.27 |
| N-155 | 1455-1485 | 42.26 | 48.73 | 9.01 |
| N-156 | 1485-1515 | 54.46 | 31.83 | 13.71 |
| N-160 | 1565-1592 | 41.68 | 24.68 | 33.64 |
| N-166 | 1675-1775 | 82.33 | 3.59 | 14.08 |
| N-168 | 1835-1842 | 35.85 | 35.62 | 28.53 |
| N-184 | 1925-1945 | 6.20 | 70.70 | 23.10 |
| N-197 | 2098-2109 | 48.23 | 22.78 | 28.99 |
| N-225 | 2364-2381 | 41.09 | 18.85 | 40.06 |
| N-226 | 2381-2400 | 5.36 | 33.83 | 60.81 |
| N-227 | 2400-2422 | 4.73 | 55.26 | 40.01 |
| N-229 | 2444-2463 | 5.28 | 82.52 | 12.2 |
| N-249 | 2768-2781 | 1.95 | 37.25 | 60.80 |

* For graphical presentation see Fig 2.17.

Table B.1 continued:

| Lab No. | Depth (cm) | Sand (%) | Silt (%) | Clay (%) |
|---------|------------|----------|----------|----------|
| N-268 | 3021-3045 | 11.75 | 28.43 | 59.82 |
| N-288 | 3278-3290 | 3.35 | 39.98 | 56.67 |
| N-350 | 4106-4120 | 12.73 | 23.31 | 63.96 |
| N-378 | 4501-4511 | 11.61 | 32.27 | 56.12 |
| N-397 | 4781-4790 | 30.21 | 30.33 | 39.46 |
| N-411 | 4986-4995 | 17.96 | 29.1 | 52.94 |
| N-412 | 4995-5000 | 18.46 | 35.75 | 45.79 |
| N-424 | 5465-5485 | 53.91 | 6.89 | 39.20 |

Table B.2 Results of clay mineral analyses* on samples from Nal Sarovar core.

| Lab No. | Depth (cm) | Smectite (%) | Kaolinite (%) | Chlorite (%) | Illite (%) |
|---------|------------|--------------|---------------|--------------|------------|
| N-4 | 10-20 | 12.96 | 7.79 | 5.20 | 74.05 |
| N-34 | 86.5-90 | 7.57 | 8.56 | 4.47 | 79.40 |
| N-75 | 195-207 | 8.29 | 7.2 | 4.86 | 79.56 |
| N-83 | 280-290 | 8.52 | 5.78 | 4.62 | 81.08 |
| N-100 | 393-417 | 12.09 | 7.46 | 3.74 | 76.71 |
| N-105 | 477-522 | 52.09 | 9.71 | 8.49 | 29.71 |
| N-113 | 600-610 | 55.76 | 8.85 | 2.21 | 33.18 |
| N-118 | 685-702 | 53.04 | 10.33 | 6.19 | 30.44 |
| N-124 | 775-790 | 42.59 | 14.68 | 4.77 | 37.96 |
| N-126 | 887-907 | 67.15 | 9.66 | 4.35 | 18.84 |
| N-132 | 987-1000 | 68.40 | 5.10 | 1.70 | 24.80 |
| N-136 | 1065-1102 | 70.30 | 5.91 | 1.97 | 21.82 |
| N-147 | 1285-1299 | 74.78 | 5.22 | 2.61 | 17.39 |
| N-148 | 1299-1324 | 62.14 | 6.19 | 3.10 | 28.57 |
| N-155 | 1455-1485 | 64.92 | 4.16 | 2.61 | 28.51 |
| N-156 | 1485-1515 | 51.43 | 7.27 | 7.01 | 34.29 |
| N-227 | 2400-2422 | 52.89 | 5.82 | 5.36 | 35.93 |
| N-229 | 2444-2463 | 97.53 | 1.19 | 0.44 | 0.84 |
| N-249 | 2768-2781 | 97.96 | 0.57 | 0.45 | 1.02 |
| N-268 | 3021-3045 | 76.47 | 7.86 | 3.89 | 11.76 |
| N-288 | 3278-3290 | 98.23 | 1.18 | 0.59 | - |
| N-350 | 4106-4120 | 98.69 | 0.66 | 0.65 | - |
| N-378 | 4501-4511 | 94.87 | 3.44 | 1.69 | - |
| N-412 | 4995-5000 | 97.45 | 1.82 | 0.73 | - |
| SPM** | - | 29.82 | 21.59 | 6.48 | 42.11 |

* For graphical presentation see Fig 2.18.

** Suspended sediment from flood water of Sabarmati river at Ahmedabad (Sept. 1994).